

Flow Energy, Time, and Evolution of Dynamic Fluvial Systems: Implications for Stabilization and Restoration of Alluvial Streams

(...A Review...)

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Introductory Points

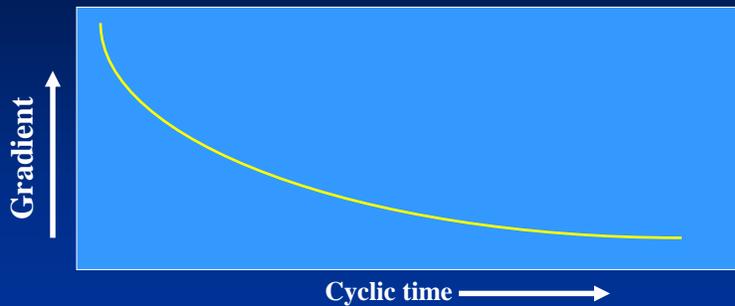
- Alluvial streams are open systems that dynamically adjust to variations in flow energy and sediment supply.
- Streams adjust their morphology to imbalances between available force and sediment supply as a function of the resistance of the boundary sediments to hydraulic and geotechnical forces.
- Thus, two (reference?) channels of similar morphology disturbed by an identical perturbation can attain different equilibrium morphologies
- Also, diverse streams subject to diverse perturbations can respond similarly.



We Restore Unstable Channels (usually in unstable systems)

- Stream restoration is conducted on unstable channels (systems) where regime relations are not valid and bankfull dimensions are dynamic or even indefinable
- Because unstable streams are dynamic, it is essential to know where in the adjustment cycle the stream is so that upstream sediment supply can be accounted for
- Although adjustments make alluvial streams particularly dynamic, response is not chaotic but proceeds in a systematic and predictable way
- Therefore, one must account for the spatial and temporal aspects of adjustment when designing a restoration project

Compression of Time Scales



Compression of time scales following large-scale disturbances: “natural” or anthropogenic

1,000,000 years = 25 - 100 years

Minimization of Energy Dissipation

Channels adjust such that their geometry provides for a minimum rate of energy dissipation given the constraints of the upstream sediment load, roughness and boundary materials

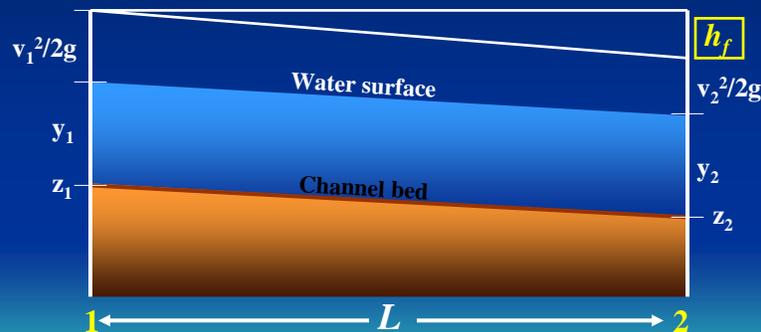
If this holds true and given the compressed time scales for adjustment, then we should be able to track this over time in disturbed, adjusting streams

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Flow Energy and Energy Dissipation

$$E = z + \gamma y + \frac{v^2}{2g}$$

$$h_f = (z_1 + \gamma y_1 + \frac{v_1^2}{2g}) - (z_2 + \gamma y_2 + \frac{v_2^2}{2g})$$



$$\text{Energy slope: } S_e = h_f / L$$

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Processes That Effect Components of Total Mechanical Energy (E)

For each parameter comprising E ,
what processes would result in a
reduction in those values?

- z : degradation
- γy : widening, aggradation
- $v^2/2g$: widening, increase in relative roughness, growth of vegetation, aggradation,

Thus, different and often opposite processes can have
the same result. Which dominate where?

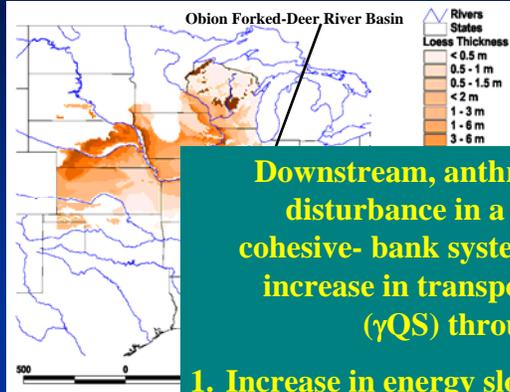
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Idealized Adjustment Trends For a given discharge (Q)



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Coastal-Plain System



Downstream, anthropogenic disturbance in a sand-bed, cohesive-bank system causing an increase in transport capacity (γQS) through:

1. Increase in energy slope
2. Increase in velocity head



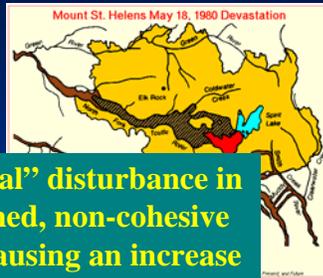
CQ

Adjustment Processes



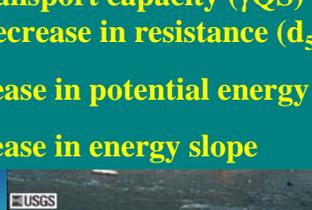
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Sub-Alpine System



Upstream “natural” disturbance in a coarse-grained, non-cohesive bank system causing an increase in transport capacity (γ_{QS}) and a decrease in resistance (d_{50}):

1. Increase in potential energy
2. Increase in energy slope



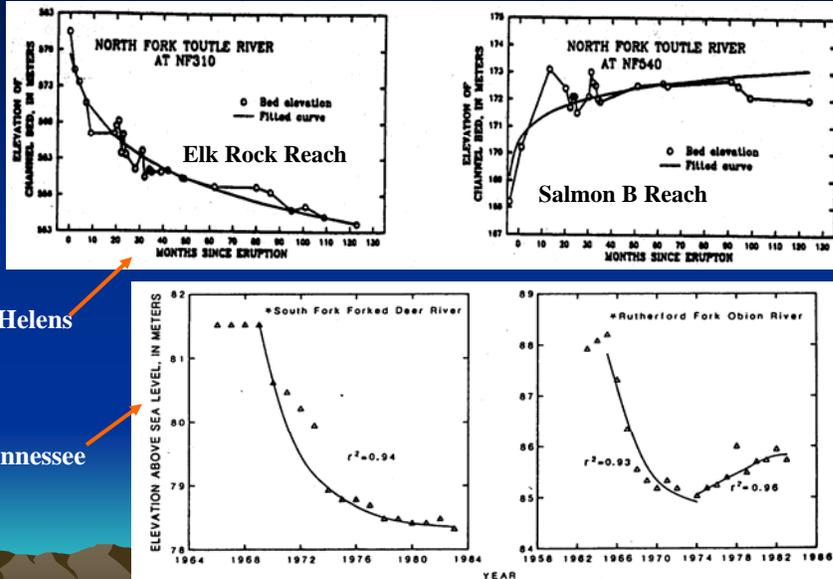
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Adjustment Processes



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Trends of Bed-Level Change

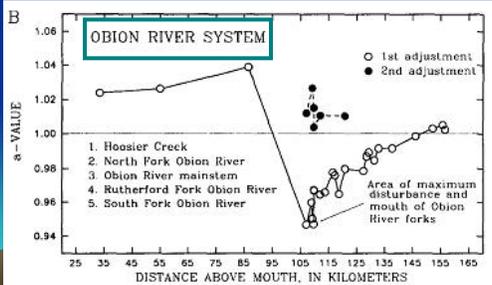
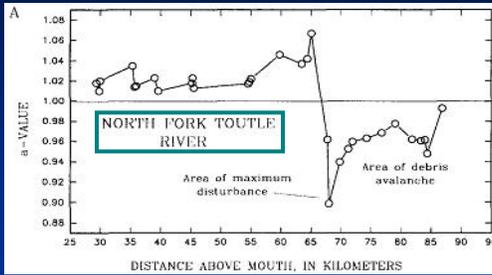


Mt. St Helens

W. Tennessee

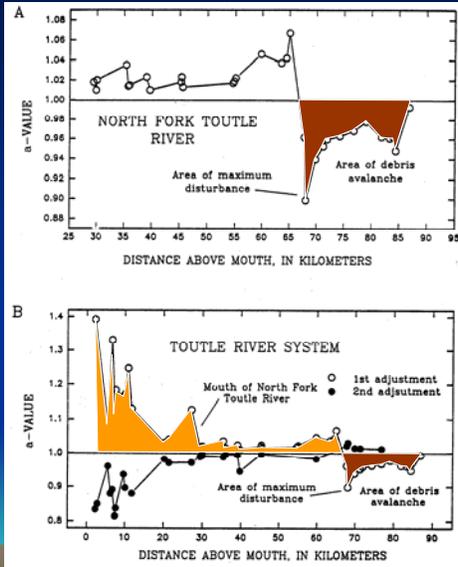


Trends of Bed-Level Change



Trends of Bed-Level Change

Coarse-grained material for aggradation derived from bank sediment.

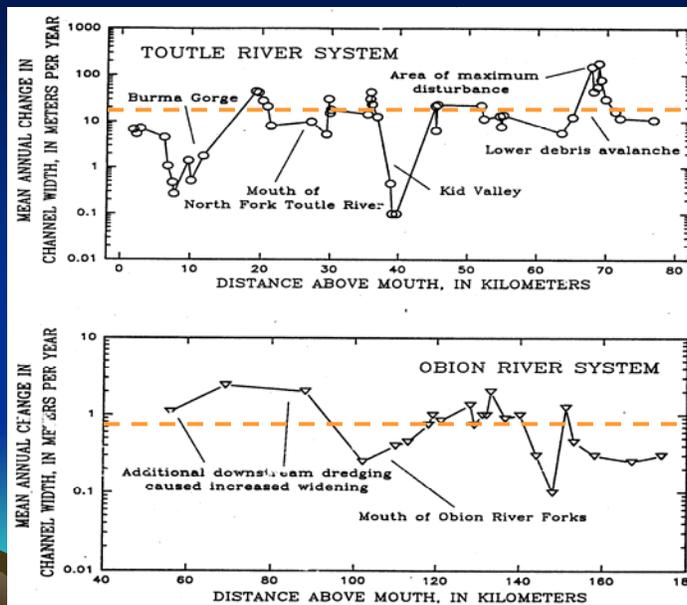


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Widening

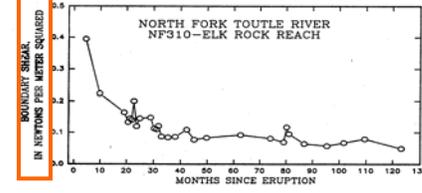
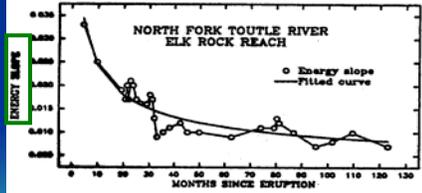
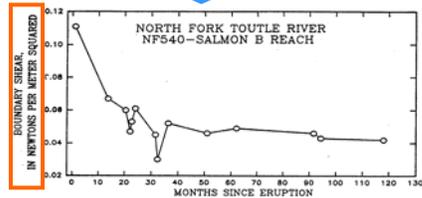
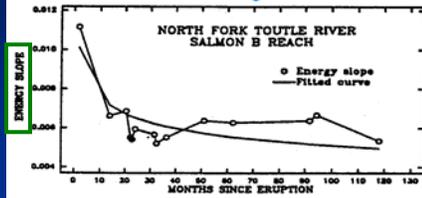
But why are they so different ?

Resistance



Adjustment by Different Processes

Aggradation and widening

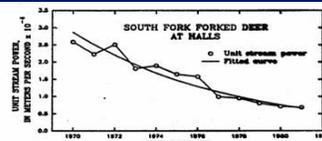
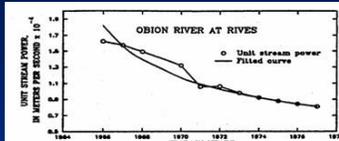


Degradation and widening

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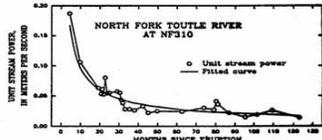
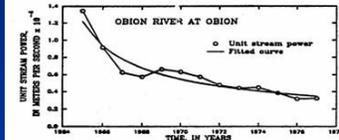
Minimization of Unit Stream Power

W. TN degrading



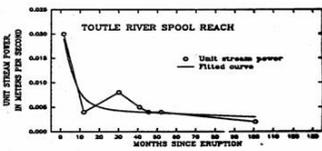
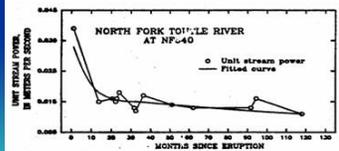
W. TN degrading

W. TN aggrading



MSH degrading

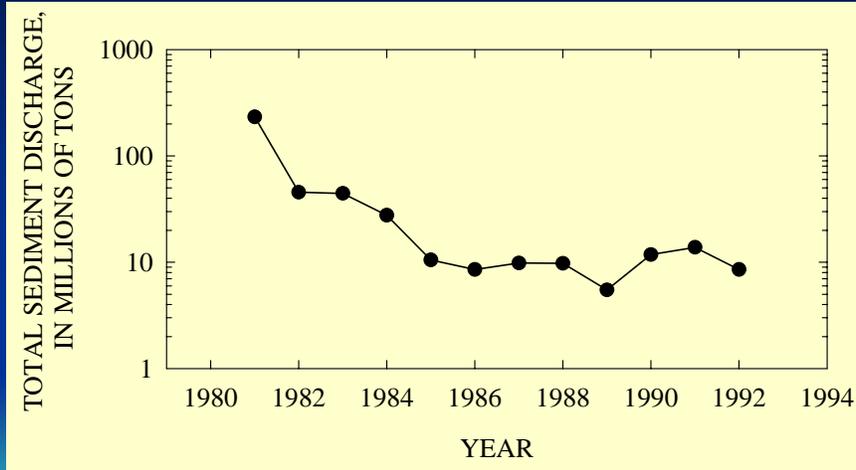
MSH aggrading



MSH aggrading

Same power adjustment, but by different processes!

Results of Energy Minimization (Sediment Discharge)



Consider this as an analog to upstream sediment supply in a recovering system

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Implications of Not Distinguishing Between Bank Materials

- Assume that $\gamma QS \propto Q_s d_{50}$ is balanced
- How does a channel respond if disturbed?
- Will the channel incise?
- Will the channel fill?
- Will the channel widen?
- Will the channel narrow?
- Will it equilibrate to the same geometry?

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Disturbing the C5 Channel

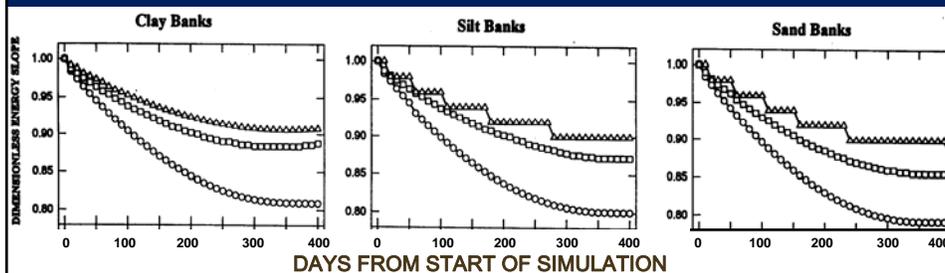
- Assume that $\gamma QS \propto Q_s d_{50}$ becomes un-balanced
- $Q_s d_{50} = 0.5$ * capacity
- Slope = 0.005
- Width/depth ratio = 13.5

| Bank material | Bed d_{50} (mm) | Bank cohesion (kPa) | Friction angle ($^{\circ}$) | Sand content (%) |
|---------------|----------------------|------------------------|----------------------------------|---------------------|
| Sand | 1.0 | 4.0 | 32.5 | 100 |
| Silt | 1.0 | 7.5 | 32.5 | 20 |
| Clay | 1.0 | 40.0 | 32.5 | 10 |

Additional cohesive strength can also be used as an analog for the root-reinforcement (mechanical) effects of vegetation on bank stability

CQ

Adjustment in Energy Slope for Different Boundary Materials and Slopes



- ▲ 0.00005 @ 0.90
- 0.0005 @ 0.87
- 0.005 @ 0.80

Energy adjustment is similar at given slope!

Do each of these channels reach equilibrium similarly?

Adjustments for Different Boundary Materials

| Gradient 0.005 m/m | |
|--------------------|------|
| Degradation | 3.51 |
| Widening | 0 |
| W_0/D_0 | 13.5 |
| W_f/D_f | 5.62 |
| τ_0 | 103 |
| τ_f | 98.7 |
| | |
| Degradation | 2.74 |
| Widening | 11.3 |
| W_0/D_0 | 13.5 |
| W_f/D_f | 8.62 |
| τ_0 | 103 |
| τ_f | 69.3 |
| | |
| Degradation | 0.35 |
| Widening | 13.1 |
| W_0/D_0 | 13.5 |
| W_f/D_f | 16.4 |
| τ_0 | 103 |
| τ_f | 64.8 |

Clay-bank channel

Silt-bank channel

Sand-bank channel

From Simon and Darby (1997)

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Response of the Channel

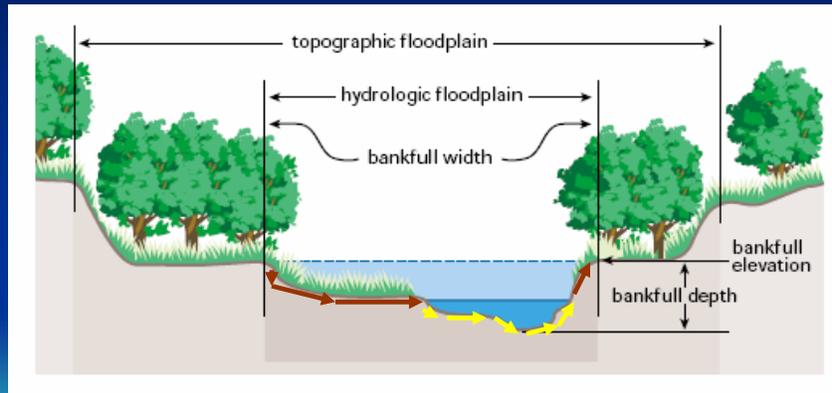
- How Does the channel respond? **It depends!**
- How much will the channel incise*? **0.4 – 3.5 m**
- How much will the channel widen*? **0 – 13 m**
- What is the stable W/D ratio*? **5.6 – 16.4**

And this is knowing the upstream sediment supply!

* For a given initial slope of 0.005 m/m

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Implications for a Form-Based System: Bed or Channel Material? (two different populations)



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Why is this Important?

- Sites may not be classified correctly: Example:

“C” channel shape: gravel bed, silt/clay banks

“C” channel shape: sand bed, sand banks

} C5

These two C5's represent very different transport regimes

- As we have seen, differences in bank materials are critical to predicting channel response and stable geometries
- Particle-size data cannot be used for incipient motion or transport analysis
- Extensive data sets collected by various agencies cannot be used for analysis of hydraulic erosion, geotechnical stability, or channel response

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Perhaps Explains Why...

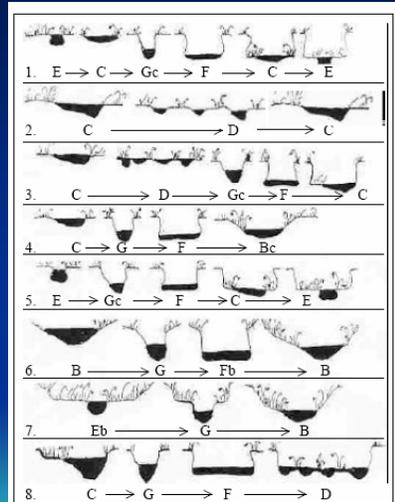


Figure 1. Various Stream Type Evolution Scenarios

From Rosgen (2001)

“The consequence of a wide range of stream channel instability can be described and **quantified** through an evolution of stream types (Figure 1).”

Rosgen (2001)

- **E to E**
- **C to C; C to Bc; C to D**
- **B to B**
- **Eb to B**

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Forcing a Form-Based System to Describe Process

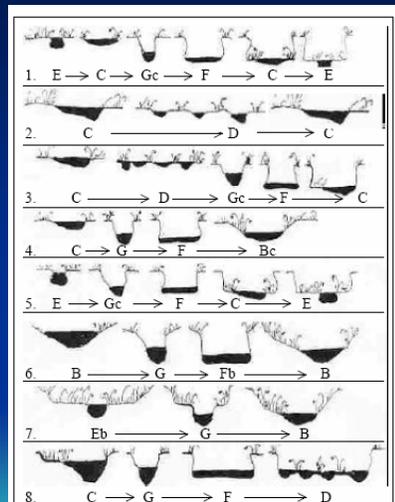


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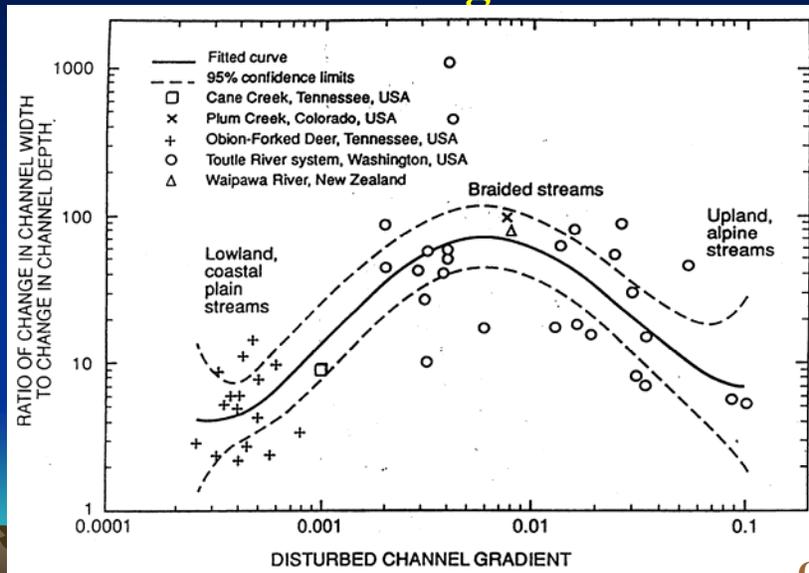
- **E to E**
- **C to C; C to Bc; C to D**
- **B to B**
- **Eb to B**

Can this be predicted a priori?

What does this mean for the reference reach-based approach?

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These Changes Occur at Different Rates and Magnitudes



Conclusions

- Restoration has to be in the context of the condition of the watershed and the channel system... **spatial and temporal aspects of the instability**
- Restoration of an unstable reach within an dynamic, unstable system will not likely be successful
- A deterministically-based approach that accounts for bed and bank processes provides a reliable means of analyzing adjustment processes and predicting stable-channel geometries in unstable systems.

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